Nitrogen turnover and N_2 : N_2O partitioning from agricultural soils – a simplified incubation assay

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Abstract

Nitrogen turnover and related denitrification losses are a major uncertainty when estimating N loss and replacement from agro-ecosystems, due to methodological constraints quantifying N₂ and laborious analytical procedures. We present a novel, simplified incubation assay that combines the ¹⁵N gas flux method with the ¹⁵N pool dilution method, to quantify denitrification losses as a function of N turnover. This assay was tested using a pasture soil from sub-tropical Australia. N-fertiliser (35 μ g g⁻¹ soil) was applied either as a single $(NH_4^{15}NO_3)$ or double $({}^{15}NH_4^{15}NO_3)$ labelled treatment at 10 atom %, with a third treatment $(NH_4^{15}NO_3)$ at 60 atom % to quantify N₂ emissions. Gross rates of N mineralisation, nitrification and related N₂ and N₂O emissions were measured during 48 hours of incubation at 80% WFPS. Gross N production and gross N consumption was consistent with the directly measured N pool sizes, with denitrification losses (N_2+N_2O) at 7.0 ± 1.4 µg N g⁻¹ soil accounting for 62% of the calculated NO₃⁻ consumption. N turnover was dominated by mineralisation and nitrification, increasing the NO_3 pool by a factor of 3. High NO_3 concentrations shifted the N₂:N₂O ratio towards N₂O, with 60 % of denitrification losses emitted as N₂O. More than 25% of the applied ¹⁵N fertiliser was lost via denitrification, showing the significance of denitrification as a major pathway of N loss from agro-ecosystems. The simplified incubation assay proved to be an efficient tool to quantify N pools and emissions, and as such is an effective method to establish comprehensive datasets of denitrification losses linked to N turnover from agro-ecosystems.

Key Words Denitrification; Dinitrogen; Nitrous oxide; ¹⁵N pool dilution; Mineralisation; Nitrification

Introduction

Denitrification losses in the form of dinitrogen (N₂) and nitrous oxide (N₂O) from agricultural soils are a major uncertainty when estimating nitrogen (N) loss and replacement, reflecting the difficulties of measuring N₂ against a high atmospheric background and laborious analytical procedures involved. Denitrification losses are tightly coupled to N-turnover in soil, where plant uptake, microbial N immobilisation, dissimilatory nitrate reduction (DNRA) and denitrification compete for the mineral N supplied into the NO₃⁻ pool. Net rates of N turnover show only the change of the mineral N pools over time, e.g. the increase or decrease of the ammonia (NH₄⁺) and nitrate (NO₃⁻) pool. In contrast, gross N production and consumption rates quantify the actual influx of N into and efflux from the respective mineral N pool, revealing the relation between production and consumption of NH₄⁺ and NO₃⁻. This is significant for both magnitude and N₂:N₂O partitioning of denitrification losses: N turnover controls the amount of NO₃⁻ available for denitrification; and high NO₃⁻ levels shift the N₂:N₂O ratio towards N₂O (Giles et al., 2012).

Denitrification is a major pathway of N-loss from upland soils, however the magnitude of N_2 and N_2O emissions and related gross rates of N turnover are still unknown for a wide range of agroecosystems (Butterbach-Bahl et al., 2013). The ¹⁵N gas flux method enables the direct quantification of N_2 and N_2O emissions after ¹⁵N fertiliser addition (Mulvaney, 1984). The ¹⁵N pool dilution method (Kirkham and Bartholomew, 1954) is based on the isotopic dilution of the ¹⁵N labelled N pool and quantifies gross rates of N turnover.

Combining both methods allows to address major uncertainties of the N-cycle: gross rates of mineralisation and nitrification, the fraction of N_2O derived from nitrification as a function of gross nitrification and the magnitude of denitrification including $N_2:N_2O$ partitioning. The objective of this study is to develop and test an efficient method to obtain comprehensive datasets on N cycling, combining the ¹⁵N gas flux method with the ¹⁵N pool dilution method in a simplified soil incubation assay.

Material and Methods

Incubation

Bulk soil samples (0-10cm) were collected from an intensively managed pasture in subtropical Casino, New South Wales, Australia. The soil is a heavy clay Vertosol with a pH of 6.6. Soil samples were air dried, sieved to < 4 mm and stored in a cold room. Prior to incubation, soil water content was determined gravimetrically after drying the soil at 105°C for 24 hours. Soil microcosms were established in centrifuge

tubes (50 ml) using the exact equivalent of 8 g oven dry soil, with 4 replicates per fertiliser treatment. NH_4NO_3 equivalent to 35 µg g⁻¹ soil was applied to each replicate (1ml solution 8g⁻¹ soil). The NH_4NO_3 was either single ($NH_4^{15}NO_3^{-}$) labelled (a) or double ($^{15}NH_4^{15}NO_3$) labelled (b) at 10 atom %. A second single labelled treatment ($NH_4^{15}NO_3^{-}$) at 60 atom % (c) was used to quantify N_2 emissions. Additional water was applied to achieve the targeted water filled pore space (WFPS) of 80% at an adjusted soil bulk density of 1 (g cm⁻³). Water and fertiliser solution were applied dropwise on two layers of 4 g of soil each to ensure homogenous ¹⁵N labelling. After fertilisation, centrifuge tubes were closed with Suba-seals (Sigma Aldrich). Soil microcosms were kept closed in an incubator at a constant temperature of 25°C.

Gas samples

Ambient air samples (n=4) were taken daily before closing the centrifuge tubes, to quantify ambient N₂O concentrations. Specific background samples were taken above the respective soil microcosms treated with $NH_4^{15}NO_3$ at 60 atom % (c) for $^{15}N_2$ analysis before closing the tubes. The entire headspace atmosphere was sampled 24 and 48 h after closure using a gas-tight syringe. After gas sampling, the Suba-seals were removed for 10 minutes, allowing the headspace atmosphere to equilibrate. Gas samples were transferred into pre-evacuated 12 ml Exetainer tubes with a double wadded Teflon/silicon septa cap (Labco Ltd, Buckinghamshire, UK) and stored until N₂O analysis by gas chromatography (Shimadzu GC-2014). Gas samples were analysed for $^{15}N_2O$ and $^{15}N_2$ using an automated isotope ratio mass spectrometer (IRMS) (Sercon Limited, 20-20, UK).

Soil samples

Soil microcosms were extracted with 40 ml 2 M KCL, immediately and 48 h after fertiliser application. Four additional soil microcosms were extracted before fertiliser application to determine starting conditions. All soil extractions were conducted in the centrifuge tubes to avoid subsampling. The centrifuge tubes were shaken with a horizontal shaker (150 rpm/min) for one hour and extracts were filtered through Whatman no. 42 filter paper. Inorganic NH_4^+ -N and NO_3^- -N in the KCL extract was determined by $AQ2^+$ (SEAL Analytical WI, USA). The ¹⁵N enrichments of the NH_4^+) and NO_3^- pool were determined by the diffusion method (Stark and Hart, 1996). To account for microbial ¹⁵N assimilation and immobilization on clay surfaces, the extracted soil samples were oven-dried at 60°C, finely ground using a planetary mill and analysed for ¹⁵N by IRMS.

Fluxes of N₂ and N₂O

The flux rates of N₂O were calculated from the slope of the linear increase in gas concentration during the closure period and were corrected for temperature and air pressure. Assuming that N₂ and N₂O originate from the same source pool undergoing denitrification, the ¹⁵N enrichment of the NO₃⁻ pool is calculated based on ¹⁵N-N₂O (Stevens and Laughlin, 2001). ¹⁵N-N₂O gas analysis enables the calculation of the fraction of N₂O (*d*'D) derived from denitrification, and the ¹⁵N enrichment of the NO₃⁻ pool undergoing denitrification (*a*D). Fluxes of N₂ were calculated using *a*D and the increase of ¹⁵N-N₂ in the chamber headspace following denitrification (Mulvaney, 1984). Flux calculations are described in detail in Friedl et al. (2016).

¹⁵N pool dilution and DNRA

Gross N mineralisation and gross NH_4^+ consumption were calculated based on the isotopic dilution of the ¹⁵NH₄⁺ pool in treatment (b), with the ¹⁵N enrichment of treatment (b) corrected for ¹⁵N recovered in the NH₄⁺ pool of treatment (a). Gross nitrification and gross NO₃⁻ consumption were calculated based on the isotopic dilution of the ¹⁵NO₃⁻ pool in treatment (a), assuming nitrified NH₄⁺ at mean ¹⁵NH₄⁺ pool excess abundance (Barraclough, 1991). DNRA was calculated as described in Huygens et al. (2008).

Results

N_2 and N_2O emissions

Gaseous N emissions and net and gross rates of N turnover are shown in Figure 1. Over the 2-day incubation, $7.01 \pm 1.39 \ \mu\text{g} \text{ N g}^{-1}$ were emitted via denitrification (N₂+N₂O). The main product of denitrification was N₂O, with cumulative fluxes of $4.3 \pm 0.9 \ \mu\text{g} \text{ N} \text{ N}_2\text{O} \text{ g}^{-1}$, exceeding cumulative N₂ fluxes of $2.8 \pm 1.1 \ \mu\text{g} \text{ N} \text{ N}_2$ g⁻¹ by a factor of 1.5. Nitrification contributed $1.3 \pm 0.1 \ \mu\text{g} \text{ N} \text{ N}_2\text{O} \text{ g}^{-1}$ to overall N₂O fluxes. The method detection limit (DL) for N₂ ranged from 0.014 $\mu\text{g} \text{ g}^{-1} \text{ day}^{-1}$ with *a*D assumed at 50 atom % to 0.127 $\mu\text{g} \text{ g}^{-1} \text{ day}^{-1}$ with *a*D assumed at 20 atom %.

N-turnover

The NH_4^+ pool remained constant over the 2 day incubation, while the NO_3^- concentrations increased by a factor of 3 to $65.5 \pm 1.5 \ \mu g \ N-NO_3^- \ g^{-1}$, reflected in the respective net rates. Gross rates of mineralisation and

nitrification averaged at 55.6 ± 11.5 and $54.5 \pm 6.5 \mu g \text{ N}\text{-} g^{-1}$ over two days respectively, and showed a high variation between soil microcosms. N₂ and N₂O emissions accounted for 67% of the calculated gross NO₃⁻ consumption.

Recovery

The fertiliser recovery for the single labelled treatment $(NH_4^{15}NO_3)$ is shown in Table 1. Recoveries were calculated based on treatment (a) and (c). Immediately after fertiliser addition, 88% percent of the applied fertiliser was recovered in the NO₃⁻ pool, with less than 1 % recovered in the soil after extraction. At the end of the incubation, more than 50% of the fertiliser remained in the NO₃⁻ pool, with only minor amounts found in the NH₄⁺ pool. N₂ and N₂O emissions accounted for more than 25% of N fertiliser applied.

Table 1: ¹⁵N fertiliser recovery in % for the single labelled treatment ($NH_4^{15}NO_3$). Values are mean ±1 standard error of the mean (n=4).

N -fertiliser recovery in %	NH4 ¹⁵ NO3
after fertiliser addition	
NH4 ⁺ pool	0.01 ± 0.00
NO ₃ pool	88.64 <u>+</u> 0.96
Immobilised	0.90 ± 0.06
Sum	89.54 <u>+</u> 1.00
at the end of the incubation	
NH4 ⁺ pool	0.23 <u>+</u> 0.04
NO ₃ pool	51.15 <u>+</u> 4.53
Immobilised	6.35 <u>+</u> 0.31
N ₂ O emission	16.47 <u>+</u> 1.19
N ₂ emissions	10.34 <u>+</u> 4.10
Sum	84.5 <u>+</u> 6.2

Discussion

The simplified soil incubation assay directly quantified denitrification losses and N_2 and N_2O partitioning as a function of gross rates of N turnover. This novel method minimizes the workload in terms of the experimental setup, sampling and soil extraction, requiring only inexpensive equipment and small amounts of ¹⁵N labelled fertiliser. As such, it allows for large numbers of samples to be processed, recommending its use to establish comprehensive datasets of denitrification losses linked to N turnover.

At 80% WFPS, N turnover was dominated by mineralisation and nitrification (Figure 1), with gross rates of >45 μ g g⁻¹ over 2 days for both processes. The gross mineralisation rate reveals the high N supply from the organic N pool in subtropical pasture soils, and the difference between gross nitrification and gross NH₄⁺ consumption indicates a potential contribution of heterotrophic nitrification to NO₃⁻ production. The relationship between these specific processes is not reflected in the respective net rates, emphasising the importance of gross rates of mineralisation and nitrification to accurately capture N cycling in the soil. Gross nitrification exceeded NO₃⁻ consumption by a factor of 5, which led to a sharp increase of the NO₃⁻ concentration in the soil. Denitrification was the main NO₃⁻ consuming process, and only minor amounts of N were recycled into the NH₄⁺ pool via DNRA, with both processes accounting for 74% of gross NO₃⁻ consumption. Significantly, the aerobic process of nitrification dominated at 80% WFPS over denitrification, which is typically favored by low O₂ availability. This can be explained by the occurrence of these processes at different microsites in the soil matrix, which is further evidenced by the substantial amount of N₂O emitted via nitrification.

Denitrification losses were dominated by N₂O emissions, exceeding N₂ emissions by a factor of 1.5. This is consistent with an incubation study using a subtropical pasture soil, reporting low product ratios of denitrification (N₂/(N₂+N₂O)) in the first 3 days at 80% WFPS (Friedl et al., 2016). At 80% WFPS, the remaining oxygen (O₂) in the soil matrix is likely to inhibit the N₂O reductase, shifting the N₂/(N₂+N₂O) ratio towards N₂O. High NO₃⁻ concentrations as observed at the end of the incubation may have also inhibited the reduction of N₂O to N₂ by decreasing the C/NO₃ ratio and/or as a consequence of preferential NO₃⁻ reduction (Giles et al., 2012).

The ¹⁵N recovery in the different N pools from the single labelled treatments of 85 % of the applied fertiliser shows that the simplified incubation assay can provide reliable results for N_2 and N_2O emissions (Table 1). The high ¹⁵N enrichment of the fertiliser and a closure time of 24 hours resulted in a DL for N_2 between 0.014 and 0.127 µg g⁻¹ day⁻¹, depending on ¹⁵N enrichment of the NO₃⁻ pool. This low DL enables the detection of significantly lower N_2 fluxes than observed in this study, an important trait when obtaining datasets on denitrification over a range of agricultural soils with a different denitrification potential.



Figure 1: Nitrogen turnover rates in the soil atmosphere - system over two days of incubation (μ g N g⁻¹ soil) at 80% WFPS. Net rates show the change of the respective N pool over time. Gross rates quantify the influx into or the efflux from the respective N pool. Values are mean ± 1 standard error of the mean (n=4).

Conclusion

Combining the ¹⁵N gas flux method and the ¹⁵N pool dilution method, this novel incubation assay accurately quantified gross rates of N turnover in the soil-atmosphere system. This study confirms denitrification as a major pathway of N loss from subtropical pasture systems, and reveals the rapid N turnover in these agroecosystems. The simplified incubation assay allows large numbers of samples to be processed, and as such is an effective method to establish comprehensive datasets on N cycling in agro-ecosystems. These datasets, while improving a quantitative process understanding for denitrification, will enable biochemical models to accurately simulate the N cycle, revealing both short and long term effects of improved management practice on N use efficiency for fertilised agro-ecosystems.

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